8–7 From Micro to Macro: Approaching Mountain Formation Processes Based on Age Determinations of Minerals

— Insights to the Ou Backbone Range Obtained from Low-Temperature Thermochronology —

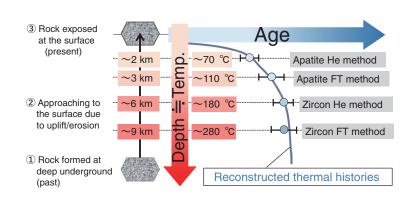


Fig.8-22 Reconstructing the time-temperature history involving uplift and cooling of a rock sample based on multi-thermochronometers After formation, a rock gradually approaches the surface due to uplift/erosion before finally becoming exposed at the surface. Multiple dating methods having various closure temperatures were applied to a single rock sample collected at the present surface. The obtained ages should vary depending on the closure temperatures/depths of the applied methods. The curve connecting each point represents a time-temperature history of the rock, i.e., a cooling history.

Sites selected for geologic disposal of high-level radioactive waste (HLW) should be stable in the long term (i.e., 10⁵ yr). At the Tono Geoscience Center (TGC), techniques for evaluating and investigating the long-term geosphere stability have been developed. In particular, estimating the uplift/erosion mechanisms plays a key role in evaluating risks of HLW approaching biospheres and simulating groundwater flow conditions related to topographic evolution, thereby contributing to the design and safety assessment of a disposal facility.

Thermochronology, a branch of radiometric dating, is a useful method to evaluate the temperature of a rock sample changing over time (i.e., the cooling history). Radiometric dating employs radioactive decay, in which radionuclides (e.g., uranium) naturally decompose into other nuclides at certain rates. At TGC, fissiontrack (FT) and (U-Th)/He (He) thermochronometers are available. These instruments can be used to determine the age of a sample by analyzing the chemical composition or crystal structure subject to micro-scale (<0.1 mm) minerals, such as apatite and zircon. Closure temperatures, below which dating methods begin to record their ages, have been determined theoretically and experimentally. The obtained age indicates when the sample was cooled to its closure temperature. Thus, an uplift/erosion history of rock can be represented by a cooling history from the hot to the cold domain (i.e., deep underground to the surface), as shown in Fig.8-22. Young ages indicate short transfer durations, implying that the rock was recently located in a warmer (i.e., deeper) location. By assuming thermal

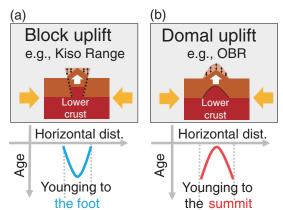


Fig.8-23 Comparing uplift mechanisms and the spatial age distributions between two mountains The spatial patterns of ages across mountains can be useful indicators to estimate the uplift mechanisms of mountains. (a) Samples closer to the foot of a mountain formed from block uplift are younger because the unstable topographies along both margins are extremely eroded. (b) Samples closer to the summit of a mountain formed by domal uplift are younger because the steepest topographies at the center are selectively eroded.

gradients, uplift/erosional histories (i.e., how quickly the rock was moved to the surface) can be reconstructed. Furthermore, a large sampling of ages across a mountain can provide a macroscale (i.e., a few to a dozen km) uplift/erosion pattern over the entire mountain.

The Tohoku district has been affected by the E–W compression for three million years due to the subduction of the Pacific Plate at the Japan Trench, which triggered the 2011 off the Pacific coast of Tohoku Earthquake. This compression has also caused the uplift of the Ou Backbone Range (OBR) along the central axis of Tohoku; however, the mechanism of uplift has not yet been clarified. Therefore, we obtained FT and He thermochronometric data at densities higher than previous studies. Our results suggested a domal uplift model of the OBR, which differs from other Japanese mountain ranges, such as the Kiso Range, as summarized in Fig.8-23. The Kiso Range is mainly uplifted by reverse faults at the foot, whereas the OBR is uplifted by the hot zone underneath it. We will continue to investigate whether such domal uplift commonly occurs in other subduction zones in Japan and overseas.

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